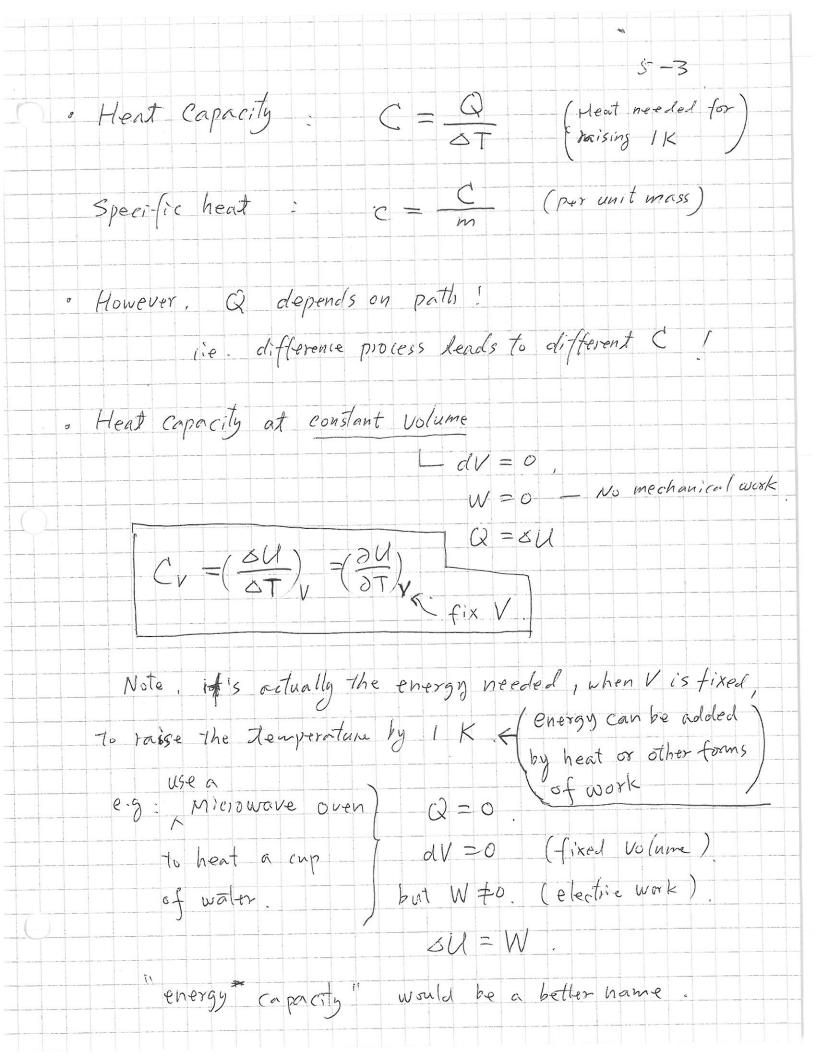


$$f V dP = -(z+f) P dV$$

$$\frac{dP}{P} = -(\frac{z+f}{f}) \frac{dV}{V}$$

$$\frac{dP}{P} = -(\frac{z+f}{f}) \frac{dV}{V}$$

$$\frac{P}{P} = -(\frac$$



For an ideal gas,
$$U = \frac{1}{2} N f R T$$
.

$$=\frac{1}{2}fnRT.$$

$$C_{V} = \left(\frac{\partial \mathcal{U}}{\partial T}\right)_{V} = \frac{d\mathcal{U}}{dT} = \frac{1}{2}Nfk = \frac{1}{2}fnR$$
(Graph on P. 30)

• Heat capacity at constant pressure
$$dQ = dU - etW$$

$$= dU + PdV$$

$$C_{P} = \left(\frac{dQ}{dT}\right)_{fix \cdot P} = \left(\frac{\partial U}{\partial T}\right)_{P} + P\left(\frac{\partial V}{\partial T}\right)_{P}$$

• For an ideal gas:
$$U = \frac{f}{2} nRT$$
, $V = \frac{nRT}{P}$.

$$\left(\frac{\partial U}{\partial T}\right)_{P} = \frac{dU}{dT} = \frac{f}{2}nR = Cv.$$

$$\left(\frac{\partial V}{\partial T}\right)_{P} = \frac{nR}{P}, \qquad \left(\frac{\partial V}{\partial T}\right)_{P} = nR$$

$$c^{*}$$
 $C_{p} = C_{v} + nR = (\frac{f}{2} + 1)nR$.

$$\frac{CP}{CV} = \left(\frac{f}{2} + 1\right) \cdot nR = \frac{f+2}{f} = \gamma$$

adiabatiz exponent

· For solids, (DV) usually is small. Cp = Cv. Since f = 6 . (Vibrations $U = \frac{f}{2} NkT = 3NkT = 3NRT$ $C_P \approx C_V = \left(\frac{\partial U}{\partial T}\right) = 3nR$ Graphs on page. 30

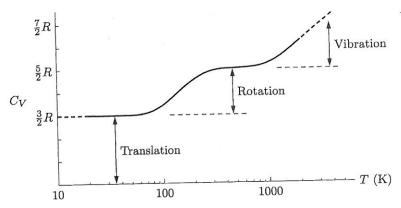


Figure 1.13. Heat capacity at constant volume of one mole of hydrogen (H₂) gas. Note that the temperature scale is logarithmic. Below about 100 K only the three translational degrees of freedom are active. Around room temperature the two rotational degrees of freedom are active as well. Above 1000 K the two vibrational degrees of freedom also become active. At atmospheric pressure, hydrogen liquefies at 20 K and begins to dissociate at about 2000 K. Data from Woolley et al. (1948).

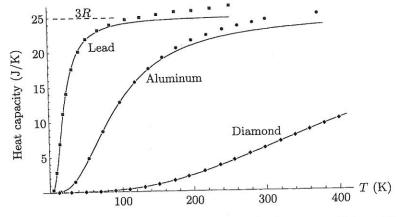
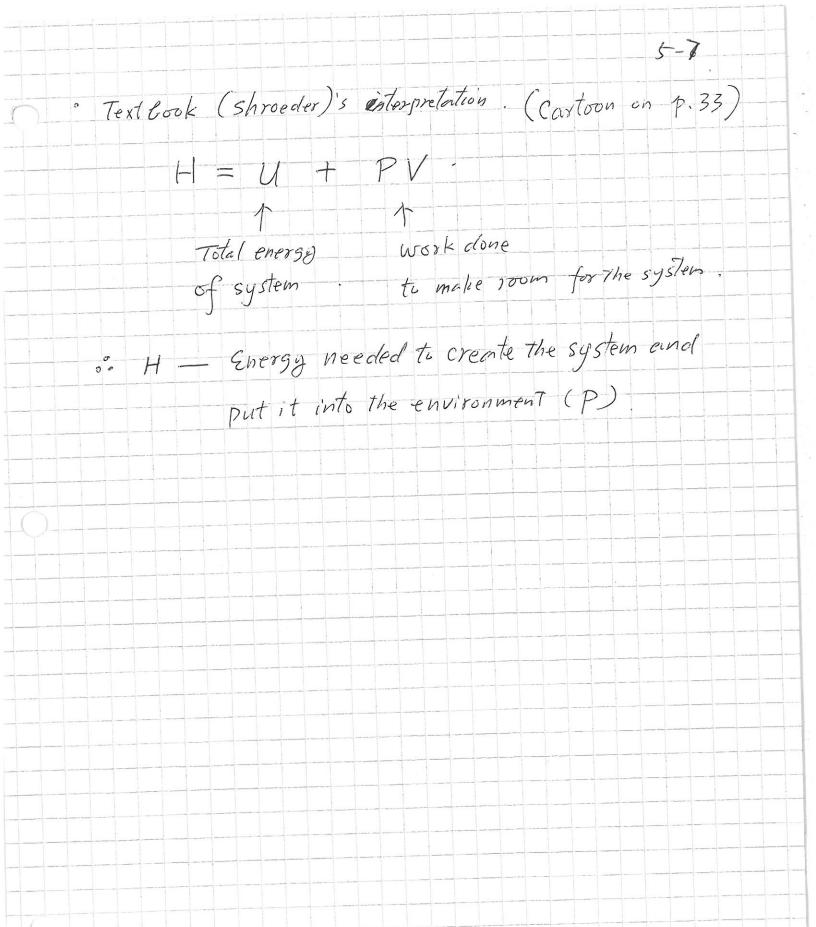


Figure 1.14. Measured heat capacities at constant pressure (data points) for one mole each of three different elemental solids. The solid curves show the heat capacity at constant volume predicted by the model used in Section 7.5, with the horizontal scale chosen to best fit the data for each substance. At sufficiently high temperatures, C_V for each material approaches the value 3R predicted by the equipartition theorem. The discrepancies between the data and the solid curves at high T are mostly due to the differences between C_P and C_V . At T=0 all degrees of freedom are frozen out, so both C_P and C_V go to zero. Data from Y. S. Touloukian, ed., Thermophysical Properties of Matter (Plenum, New York, 1970).

Therefore, $C_P = C_V + Nk = C_V + nR$ (ideal gas). (1.48)

In other words, for each mole of an ideal gas, the heat capacity at constant pressure exceeds the heat capacity at constant volume by R, the gas constant. Oddly, the

J	Latent Heat.
	- Required for phase transitions
	(without changing temperature)
	L = Q m .
Dat	ta in handbooks: usually at P= 1 atm = constant.
	e.g. melting ice: $L = 333 \text{J/g}$.
	boiting water: (= 2260 5/3 - (Un pourije)
0	$\mathcal{L}_{n+halpy}$ $H = \mathcal{U} + \mathcal{P} \mathcal{V}$
	A state quantily, like energy, that is convenient
	when dealing with heat in a constant-pressure process
	Very common, basically anyth
	occurs in the open air (P= Pax + con
	e.g. chemical reaction.



besides the usual constant-pressure expansion or compression. The latent heat for melting ice is 333 J/g, or 80 cal/g. The latent heat for boiling water is 2260 J/g, or 540 cal/g. (To get a feel for these numbers, recall that raising the temperature of water from 0°C to 100°C requires 100 cal/g.)

Problem 1.47. Your 200-g cup of tea is boiling-hot. About how much ice should you add to bring it down to a comfortable sipping temperature of 65°C? (Assume that the ice is initially at -15°C. The specific heat capacity of ice is 0.5 cal/g·°C.)

Problem 1.48. When spring finally arrives in the mountains, the snow pack may be two meters deep, composed of 50% ice and 50% air. Direct sunlight provides about 1000 watts/m² to earth's surface, but the snow might reflect 90% of this energy. Estimate how many weeks the snow pack should last, if direct solar radiation is the only source of energy.

Enthalpy

Constant-pressure processes occur quite often, both in the natural world and in the laboratory. Keeping track of the compression-expansion work done during these processes gets to be a pain after a while, but there is a convenient trick that makes it a bit easier. Instead of always talking about the energy content of a system, we can agree to always add in the work needed to make room for it (under a constant pressure, usually 1 atm). This work is PV, the pressure of the environment times the total volume of the system (that is, the total space you would need to clear out to make room for it). Adding PV onto the energy gives a quantity called the enthalpy, denoted H:

$$H \equiv U + PV. \tag{1.51}$$

This is the total energy you would have to come up with, to create the system out of nothing and put it into this environment (see Figure 1.15). Or, put another way, if you could somehow annihilate the system, the energy you could extract is not just U, but also the work (PV) done by the atmosphere as it collapses to fill the vacuum left behind.



Figure 1.15. To create a rabbit out of nothing and place it on the table, the magician must summon up not only the energy U of the rabbit, but also some additional energy, equal to PV, to push the atmosphere out of the way to make room. The *total* energy required is the enthalpy, H = U + PV.