

## Lecture 12 - The Earth's rotation

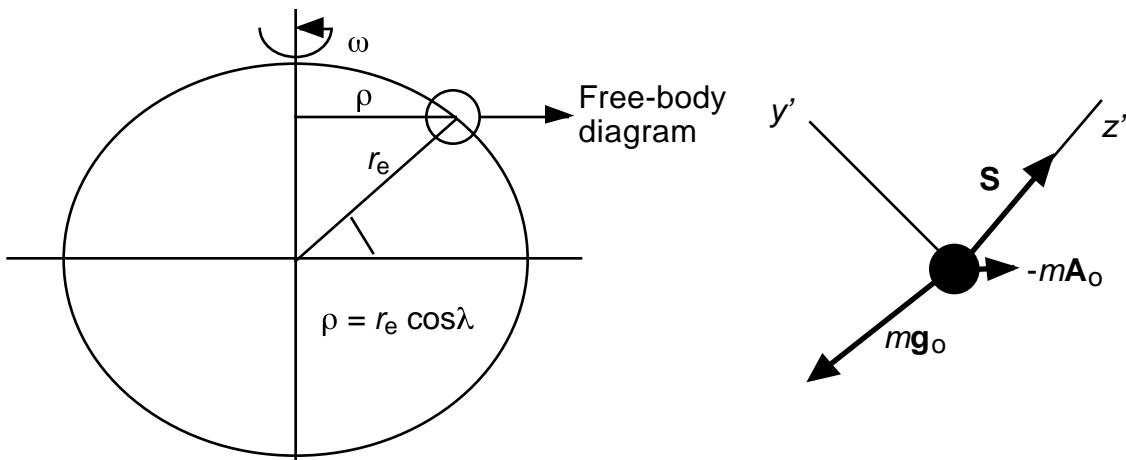
Text: Fowles and Cassiday, Chap. 5

Demo: basketball for the earth, cube for Cartesian frame

A coordinate system sitting with fixed orientation on the Earth's surface (i.e., having the  $y$ -axis pointing towards the North Pole, the  $x$ -axis pointing east along a line of latitude and the  $z$ -axis perpendicular to the Earth's surface) is both a rotating and an accelerating reference frame. Hence, fictitious forces are required in this frame to account for an object's motion. The Earth rotates counter-clockwise as seen at North Pole, so  $\omega$  is up.

### The Plumb Line

Our first example is a simple plumb line: a point mass hung from a string. The coordinate system is illustrated below



$\lambda$  = latitude (angle)

$\rho$  = distance of plumb line from axis of rotation

$r_e$  = distance of plumb line from the centre of the Earth

We've drawn the Earth as if it has a slight bulge at the equator. Although we've drawn the tension in the plumb line  $\mathbf{S}$  as if it points along  $z'$ , in fact we don't know that is the case yet (we use  $S$  instead of  $T$ , which is reserved for period).

This is a particularly simple example because the plumb bob is stationary in the rotating frame. How do we describe the bob?

- (i) it is stationary in the rotating frame:  $\mathbf{a}' = 0$  and  $\mathbf{v}' = 0$
- (ii) it sits at the origin of the frame so  $\mathbf{r}' = 0$
- (iii) the Earth's rotational speed is constant, so  $d\omega / dt = 0$

Thus, of the expression

$$m\mathbf{a}' = m\mathbf{a} - 2m\omega\mathbf{x}\mathbf{v}' - m(d\omega/dt)\mathbf{x}\mathbf{r}' - m\omega\mathbf{x}(\omega\mathbf{x}\mathbf{r}') - m\mathbf{A}_o$$

we are left with

$$0 = m\mathbf{a} - m\mathbf{A}_o \quad \text{or just } \mathbf{F} = m\mathbf{A}_o$$

Now the physical forces  $\mathbf{F}$  acting on the bob are the true gravitational force  $m\mathbf{g}_o$  toward the centre of the Earth and the tension  $\mathbf{S}$  in the string.



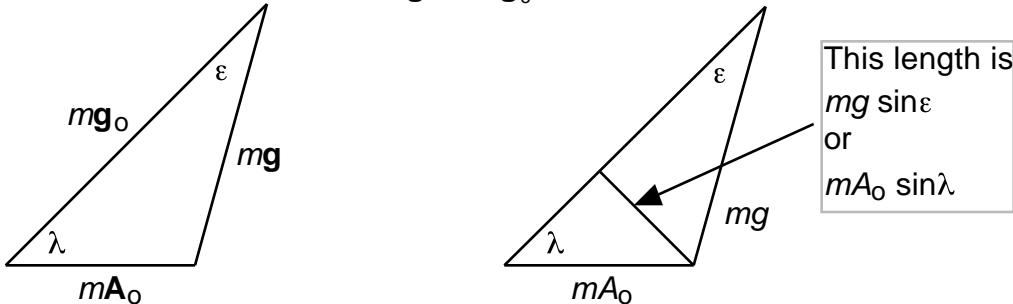
The tension balances the locally measured acceleration due to gravity  $\mathbf{g}$ , which is different from the true acceleration  $\mathbf{g}_o$ . That is

$$\mathbf{S} = -m\mathbf{g}$$

$$\Rightarrow \mathbf{F} = -m\mathbf{g} + m\mathbf{g}_o = m\mathbf{A}_o \quad (\text{vector equation})$$

or  $\mathbf{g} = \mathbf{g}_o - \mathbf{A}_o$

To evaluate the difference between  $\mathbf{g}$  and  $\mathbf{g}_o$ , we use the law of sines



$$\sin \epsilon / m\mathbf{A}_o = \sin \lambda / mg$$

$\mathbf{A}_o$  is the centripetal acceleration (and  $-\mathbf{A}_o$  is the centrifugal acceleration) given by

$$\mathbf{A}_o = \omega^2 \rho = \omega^2 r_e \cos \lambda$$

$$\Rightarrow \sin \epsilon = m\omega^2 r_e \sin \lambda \cos \lambda / mg$$

$$= \omega^2 r_e \sin 2\lambda / 2g$$

As expected,  $\epsilon$  vanishes at the equator ( $\lambda = 0$ ) and at poles ( $2\lambda = 180$  degrees). The largest value of  $\epsilon$  is at  $\lambda = 45$  degrees:

$$\omega = 2 \pi / (24 \times 3600) = 0.000073 = 7.3 \times 10^{-5} \text{ radians/sec}$$

$$r_e = 6400 \text{ km} = 6.4 \times 10^6 \text{ m}$$

$$\sin \epsilon \sim \epsilon = (7.3 \times 10^{-5})^2 \cdot 6.4 \times 10^6 \cdot 1 / (2 \times 9.8) = 0.0017 \text{ radians}$$

$$\text{or } \epsilon \sim 0.0017 \times 57.3 = 0.1 \text{ of a degree.}$$

Lastly, because the rotational motion affects both the surface of the Earth as well as the bob, then  $\mathbf{S}$  remains perpendicular to the Earth's surface. We use this fact in the

next example, and always subsume  $\mathbf{A}_o$  with  $\mathbf{g}_o$  to allow us to work with the local acceleration  $\mathbf{g}$ . At the equator,  $A_o = (7.3 \times 10^{-5})^2 \cdot 6.4 \times 10^6 = 0.034 \text{ m/s}^2$ .

### Foucault Pendulum

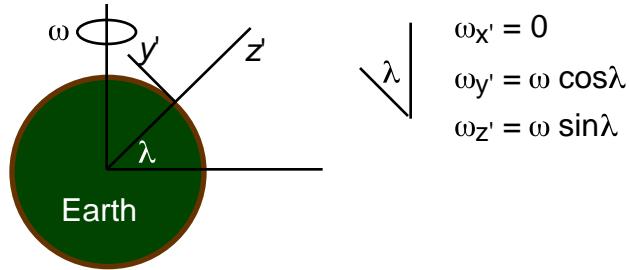
Now we take the plumb bob from the previous example and allow it to swing back and forth. The Foucault pendulum is simply a mass on a string, but the string is not restricted to move in a particular plane.

The forces acting on the pendulum bob are

$$m(d^2\mathbf{r}' / dt^2) = m\mathbf{g} + \mathbf{S} - 2m\omega\mathbf{x}\mathbf{v}',$$

where we have rolled  $\mathbf{g}_o - \mathbf{A}_o$  into the local acceleration  $\mathbf{g}$ . But because the tension  $\mathbf{S}$  changes magnitude and direction as the bob swings, it is not always true that  $\mathbf{S} = -m\mathbf{g}$ . The local centripetal acceleration term  $\omega\mathbf{x}(\omega\mathbf{x}\mathbf{r}')$  has been dropped in favour of the Coriolis force, which is much more important for this problem [The local force  $\omega\mathbf{x}(\omega\mathbf{x}\mathbf{r}')$  depends upon  $\mathbf{r}'$  which is with respect to the  $x'y'z'$  origin on the surface of the Earth, whereas  $\mathbf{A}_o$  depends on  $r_e \gg r'$ ].

Choose  $x'$ ,  $y'$  to form a plane tangent to the Earth's surface, and choose  $z'$  to therefore lie along  $\mathbf{g}$ :

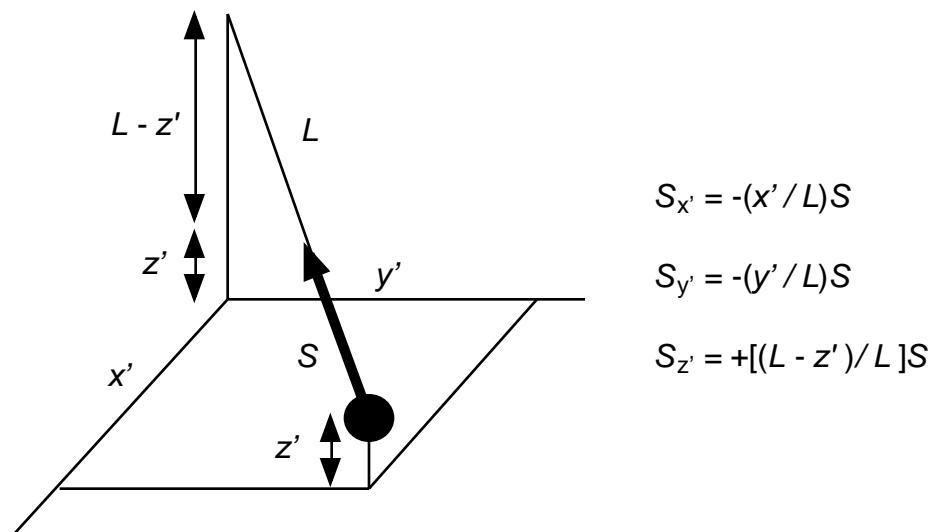


From  $\mathbf{v}' = v_x\mathbf{i}' + v_y\mathbf{j}' + v_z\mathbf{k}'$ , the Cartesian components of  $\omega\mathbf{x}\mathbf{v}'$  are

$$\begin{aligned} \omega\mathbf{x}\mathbf{v}' &= [\omega \cos\lambda (dz' / dt) - \omega \sin\lambda (dy' / dt), \\ &\quad \omega \sin\lambda (dx' / dt), \\ &\quad -\omega \cos\lambda (dx' / dt)] \end{aligned}$$

Now the components of the tension  $\mathbf{S}$  can be written in terms of the length of the string and the coordinates of the bob: comparing  $S_x / S$  with  $x' / L$ , for example, gives

$$S_x = -(x' / L)S.$$



Note that the - signs are required to obtain the correct orientations.

Returning to our equation for  $\mathbf{a}'$ , we can write out the  $x'$  and  $y'$  components as

$$m \frac{d^2x'}{dt^2} = -(x' S/L) - 2m\omega [(dz'/dt) \cos\lambda - (dy'/dt) \sin\lambda]$$

$$m \frac{d^2y'}{dt^2} = -(y' S/L) - 2m\omega (dx'/dt) \sin\lambda$$

We make two approximations in the small angle situation, where the motion is nearly horizontal:

- $dz'/dt \sim 0$  corresponding to no vertical motion
  - $S = mg$  (in magnitude) since the pendulum is almost vertical

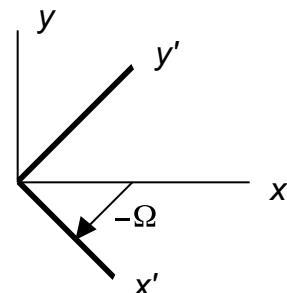
$$\Rightarrow \begin{aligned} d^2x' / dt^2 &= -(g / L) x' + 2\Omega (dy' / dt) \\ d^2y' / dt^2 &= -(g / L) y' - 2\Omega (dx' / dt) \end{aligned}$$

where  $\Omega = \omega \sin\lambda$  is the component of  $\omega$  in the  $z'$  direction, i.e., it's the local vertical component of  $\omega$ .

Somewhat like the  $q\mathbf{v}\times\mathbf{B}$  problem of charged particles in a magnetic field, the equations of motion are now coupled in the  $x'$  and  $y'$  directions. They can be uncoupled by defining yet another rotating coordinate system  $X, Y$  which rotates **clockwise** around  $z'$  with an angular frequency of  $-\Omega = -\omega \sin\lambda$

$$x' = X \cos \Omega t + Y \sin \Omega t$$

$$y' = -X \sin \Omega t + Y \cos \Omega t$$



Substituting this transformation into the first equation for  $d^2x'/dt^2$ :

$$d^2x'/dt^2 = -(g/L)x' + 2\Omega dy'/dt$$

becomes

$$d^2(X \cos\Omega t + Y \sin\Omega t)/dt^2 = -(g/L)(X \cos\Omega t + Y \sin\Omega t) + 2\Omega d(-X \sin\Omega t + Y \cos\Omega t)/dt$$

or

$$\begin{aligned} d/dt \{ dX/dt \cos\Omega t - X\Omega \sin\Omega t + dY/dt \sin\Omega t + Y\Omega \cos\Omega t \} = \\ -(g/L)(X \cos\Omega t + Y \sin\Omega t) \\ + 2\Omega \{ -dX/dt \sin\Omega t - X\Omega \cos\Omega t + dY/dt \cos\Omega t - Y\Omega \sin\Omega t \} \end{aligned}$$

Dropping terms of order  $\Omega^2$

$$(d^2X/dt^2 + gX/L) \cos\Omega t + (d^2Y/dt^2 + gY/L) \sin\Omega t = 0.$$

Each of the coefficients of  $\cos\Omega t$  or  $\sin\Omega t$  must vanish for arbitrary time, leaving just the usual simple harmonic motion equations for  $X$ ,  $Y$ :

$$\begin{aligned} d^2X/dt^2 + (g/L)X &= 0 \\ d^2Y/dt^2 + (g/L)Y &= 0 \end{aligned}$$

That these equations give the usual SHM (including the period  $2\pi\sqrt{L/g}$ ) is really not very surprising. What is important to note is that the  $XY$  coordinates rotate with an angular frequency  $\omega \sin\lambda = \Omega$  with respect to  $x'y'$ : that is, the plane of oscillation rotates with respect to a coordinate system on the Earth's surface. Using

$$\omega = 2\pi / T$$

the period of rotation  $T_{\text{foucault}}$  is then

$$T_{\text{foucault}} = T_{\text{earth}} / \sin\lambda \quad \text{where } T_{\text{earth}} = 24 \text{ hours.}$$

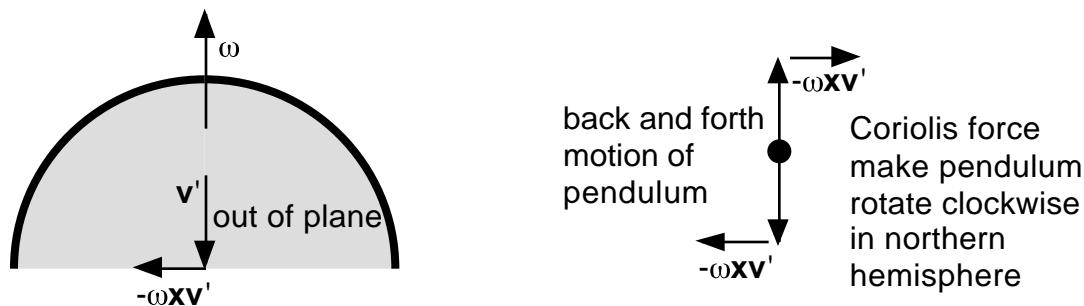
Examples:

North pole,  $\lambda = 90^\circ$ ,  $T' = 24$  hours

Equator,  $\lambda = 0$ ,  $T_{\text{foucault}} = \infty$  (no rotation)

$\lambda = 45$  degrees,  $T_{\text{foucault}} = 24 / (1/\sqrt{2}) = (\sqrt{2})(24) = 34$  hours.

The effect was first demonstrated by French physicist Jean Foucault in 1851; pendulum rotates clockwise in northern hemisphere



Effects on surface winds:

