

Lecture 12 - The Earth's rotation

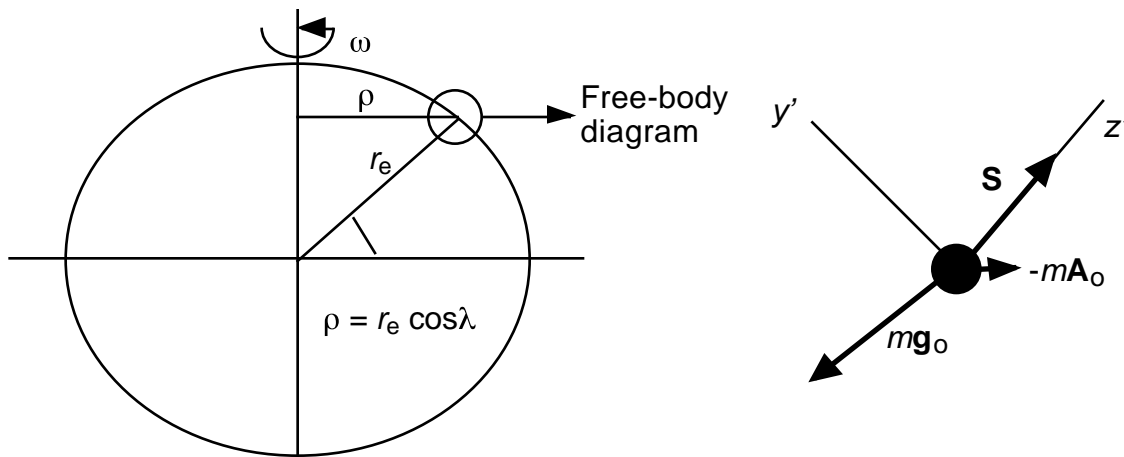
Text: Fowles and Cassiday, Chap. 5

Demo: basketball for the earth, cube for Cartesian frame

A coordinate system sitting with fixed orientation on the Earth's surface (*i.e.*, having the *y*-axis pointing towards the North Pole, the *x*-axis pointing east along a line of latitude and the *z*-axis perpendicular to the Earth's surface) is both a rotating and an accelerating reference frame. Hence, fictitious forces are required in this frame to account for an object's motion. The Earth rotates counter-clockwise as seen at North Pole, so ω is up.

The Plumb Line

Our first example is a simple plumb line: a point mass hung from a string. The coordinate system is illustrated below



λ = latitude (angle)

ρ = distance of plumb line from axis of rotation

r_e = distance of plumb line from the centre of the Earth

We've drawn the Earth as if it has a slight bulge at the equator. Although we've drawn the tension in the plumb line \mathbf{S} as if it points along z' , in fact we don't know that is the case yet (*we use S instead of T, which is reserved for period*).

This is a particularly simple example because the plumb bob is stationary in the rotating frame. How do we describe the bob?

- (i) it is stationary in the rotating frame: $\mathbf{a}' = 0$ and $\mathbf{v}' = 0$
- (ii) it sits at the origin of the frame so $\mathbf{r}' = 0$
- (iii) the Earth's rotational speed is constant, so $d\omega / dt = 0$

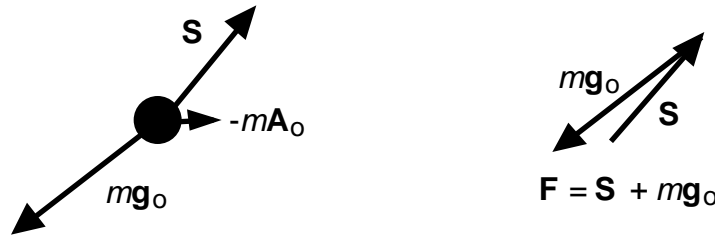
Thus, of the expression

$$m\mathbf{a}' = m\mathbf{a} - 2m\omega \mathbf{xv}' - m(d\omega / dt)\mathbf{xr}' - m\omega \mathbf{x}(\omega \mathbf{xr}') - m\mathbf{A}_0$$

we are left with

$$0 = m\mathbf{a} - m\mathbf{A}_o \quad \text{or just } \mathbf{F} = m\mathbf{A}_o$$

Now the physical forces \mathbf{F} acting on the bob are the true gravitational force $m\mathbf{g}_o$ toward the centre of the Earth and the tension \mathbf{S} in the string.



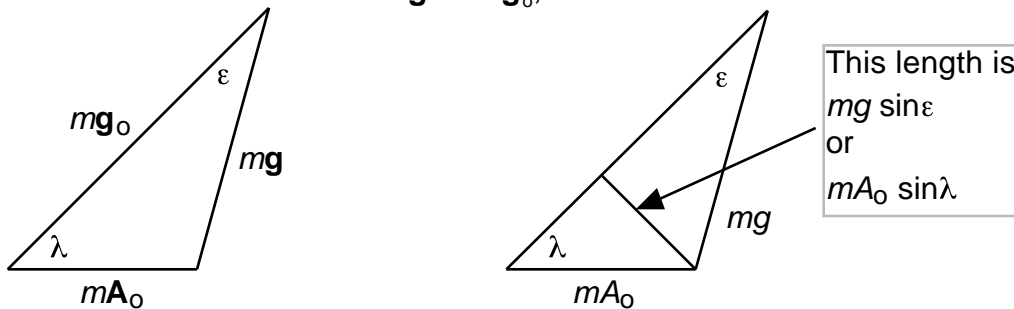
The tension balances the locally measured acceleration due to gravity \mathbf{g} , which is different from the true acceleration \mathbf{g}_o . That is

$$\mathbf{S} = -m\mathbf{g}$$

$$\Rightarrow \mathbf{F} = -m\mathbf{g} + m\mathbf{g}_o = m\mathbf{A}_o \quad (\text{vector equation})$$

$$\text{or } \mathbf{g} = \mathbf{g}_o - \mathbf{A}_o$$

To evaluate the difference between \mathbf{g} and \mathbf{g}_o , we use the law of sines



$$\sin \epsilon / mA_o = \sin \lambda / mg$$

\mathbf{A}_o is the centripetal acceleration (and $-\mathbf{A}_o$ is the centrifugal acceleration) given by

$$A_o = \omega^2 \rho = \omega^2 r_e \cos \lambda$$

$$\begin{aligned} \rightarrow \sin \epsilon &= m\omega^2 r_e \sin \lambda \cos \lambda / mg \\ &= \omega^2 r_e \sin 2\lambda / 2g \end{aligned}$$

As expected, ϵ vanishes at the equator ($\lambda = 0$) and at poles ($2\lambda = 180$ degrees). The largest value of ϵ is at $\lambda = 45$ degrees:

$$\omega = 2\pi / (24 \times 3600) = 0.000073 = 7.3 \times 10^{-5} \text{ radians/sec}$$

$$r_e = 6400 \text{ km} = 6.4 \times 10^6 \text{ m}$$

$$\sin \epsilon \sim \epsilon = (7.3 \times 10^{-5})^2 \cdot 6.4 \times 10^6 \cdot 1 / (2 \times 9.8) = 0.0017 \text{ radians}$$

or $\epsilon \sim 0.0017 \times 57.3 = 0.1$ of a degree.

Lastly, because the rotational motion affects both the surface of the Earth as well as the bob, then \mathbf{S} remains perpendicular to the Earth's surface. We use this fact in the

next example, and always subsume \mathbf{A}_o with \mathbf{g}_o to allow us to work with the local acceleration \mathbf{g} . At the equator, $A_o = (7.3 \times 10^{-5})^2 \cdot 6.4 \times 10^6 = 0.034 \text{ m/s}^2$.

Foucault Pendulum

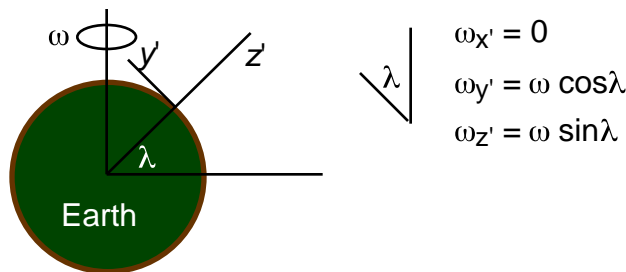
Now we take the plumb bob from the previous example and allow it to swing back and forth. The Foucault pendulum is simply a mass on a string, but the string is not restricted to move in a particular plane.

The forces acting on the pendulum bob are

$$m (d^2\mathbf{r}' / dt^2) = m\mathbf{g} + \mathbf{S} - 2m\boldsymbol{\omega}\mathbf{x}\mathbf{v}',$$

where we have rolled $\mathbf{g}_o - \mathbf{A}_o$ into the local acceleration \mathbf{g} . But because the tension \mathbf{S} changes magnitude and direction as the bob swings, it is not always true that $\mathbf{S} = -m\mathbf{g}$. The local centripetal acceleration term $\boldsymbol{\omega}\mathbf{x}(\boldsymbol{\omega}\mathbf{x}\mathbf{r}')$ has been dropped in favour of the Coriolis force, which is much more important for this problem [The local force $\boldsymbol{\omega}\mathbf{x}(\boldsymbol{\omega}\mathbf{x}\mathbf{r}')$ depends upon \mathbf{r}' which is with respect to the $x'y'z'$ origin on the surface of the Earth, whereas \mathbf{A}_o depends on $r_e \gg r'$].

Choose x', y' to form a plane tangent to the Earth's surface, and choose z' to therefore lie along \mathbf{g} :

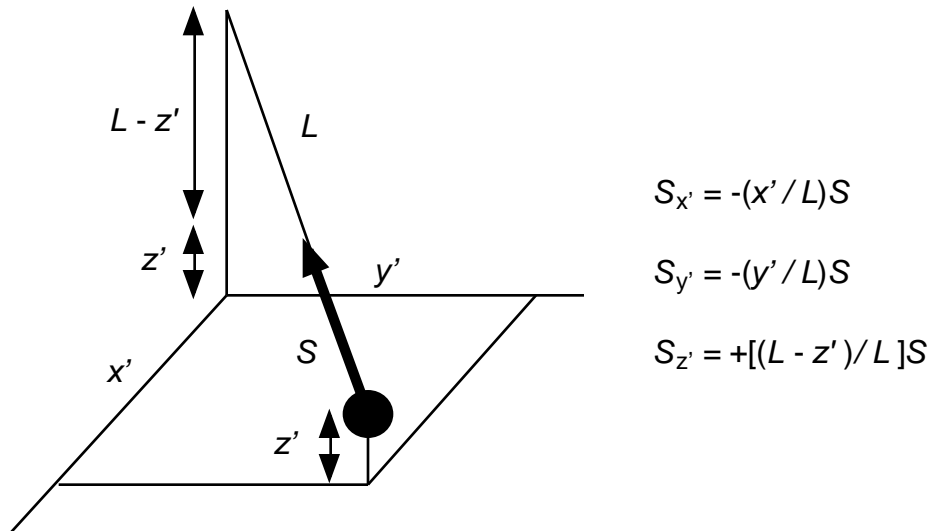


From $\mathbf{v}' = v_x\mathbf{i}' + v_y\mathbf{j}' + v_z\mathbf{k}'$, the Cartesian components of $\boldsymbol{\omega}\mathbf{x}\mathbf{v}'$ are

$$\boldsymbol{\omega}\mathbf{x}\mathbf{v}' = [\omega \cos\lambda (dz' / dt) - \omega \sin\lambda (dy' / dt), \quad \omega \sin\lambda (dx' / dt), \quad -\omega \cos\lambda (dx' / dt)]$$

Now the components of the tension \mathbf{S} can be written in terms of the length of the string and the coordinates of the bob: comparing S_x / S with x' / L , for example, gives

$$S_x = -(x' / L)S.$$



Note that the - signs are required to obtain the correct orientations.

Returning to our equation for \mathbf{a}' , we can write out the x' and y' components as

$$m (d^2x' / dt^2) = -(x' S / L) - 2m\omega [(dz' / dt) \cos\lambda - (dy' / dt) \sin\lambda]$$

$$m (d^2y' / dt^2) = -(y' S / L) - 2m\omega (dx' / dt) \sin\lambda$$

We make two approximations in the small angle situation, where the motion is nearly horizontal:

- $dz' / dt \sim 0$ corresponding to no vertical motion
- $S = mg$ (in magnitude) since the pendulum is almost vertical

$$\Rightarrow \quad d^2x' / dt^2 = -(g / L) x' + 2\Omega (dy' / dt)$$

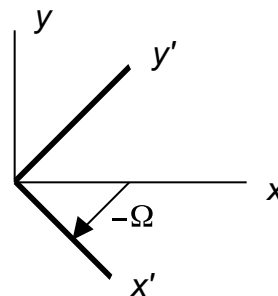
$$d^2y' / dt^2 = -(g / L) y' - 2\Omega (dx' / dt)$$

where $\Omega = \omega \sin\lambda$ is the component of ω in the z' direction, *i.e.*, it's the local vertical component of ω .

Somewhat like the $q\mathbf{v}\times\mathbf{B}$ problem of charged particles in a magnetic field, the equations of motion are now coupled in the x' and y' directions. They can be uncoupled by defining yet another rotating coordinate system X, Y which rotates **clockwise** around z' with an angular frequency of $-\Omega = -\omega \sin\lambda$

$$x' = X \cos\Omega t + Y \sin\Omega t$$

$$y' = -X \sin\Omega t + Y \cos\Omega t$$



Substituting this transformation into the first equation for d^2x'/dt^2 :

$$d^2x'/dt^2 = -(g/L)x' + 2\Omega dy'/dt$$

becomes

$$d^2(X \cos \Omega t + Y \sin \Omega t)/dt^2 = -(g/L)(X \cos \Omega t + Y \sin \Omega t) + 2\Omega d(-X \sin \Omega t + Y \cos \Omega t)/dt$$

or

$$d/dt\{dX/dt \cos \Omega t - X\Omega \sin \Omega t + dY/dt \sin \Omega t + Y\Omega \cos \Omega t\} = -(g/L)(X \cos \Omega t + Y \sin \Omega t) + 2\Omega\{-dX/dt \sin \Omega t - X\Omega \cos \Omega t + dY/dt \cos \Omega t - Y\Omega \sin \Omega t\}$$

Dropping terms of order Ω^2

$$(d^2X/dt^2 + gX/L) \cos \Omega t + (d^2Y/dt^2 + gY/L) \sin \Omega t = 0.$$

Each of the coefficients of $\cos \Omega t$ or $\sin \Omega t$ must vanish for arbitrary time, leaving just the usual simple harmonic motion equations for X , Y :

$$d^2X/dt^2 + (g/L)X = 0$$

$$d^2Y/dt^2 + (g/L)Y = 0$$

That these equations give the usual SHM (including the period $2\pi\sqrt{L/g}$) is really not very surprising. What is important to note is that the XY coordinates rotate with an angular frequency $\omega \sin \lambda = \Omega$ with respect to $x'y'$: that is, the plane of oscillation rotates with respect to a coordinate system on the Earth's surface. Using

$$\omega = 2\pi/T$$

the period of rotation T_{foucault} is then

$$T_{\text{foucault}} = T_{\text{earth}} / \sin \lambda \quad \text{where } T_{\text{earth}} = 24 \text{ hours.}$$

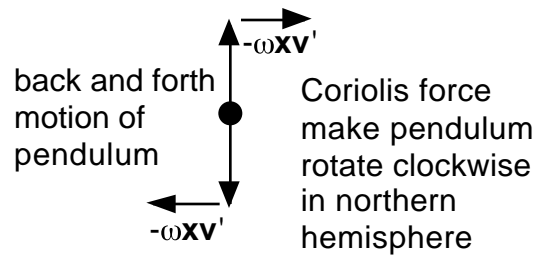
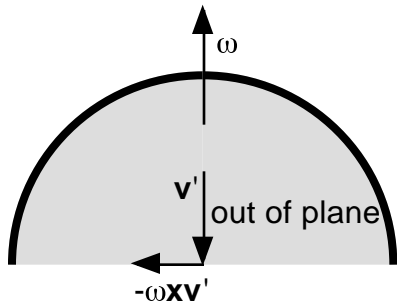
Examples:

$$\text{North pole, } \lambda = \pi/2, T' = 24 \text{ hours}$$

$$\text{Equator, } \lambda = 0, T_{\text{foucault}} = \infty \text{ (no rotation)}$$

$$\lambda = 45 \text{ degrees, } T_{\text{foucault}} = 24 / (\sin 45^\circ) = (2\sqrt{2})(24) = 34 \text{ hours.}$$

The effect was first demonstrated by French physicist Jean Foucault in 1851; pendulum rotates clockwise in northern hemisphere



Effects on surface winds:

